

# Quantifying seafloor dynamics of organic matter in the Barents Sea shelf sediments

Felipe S. Freitas<sup>1</sup>, K.R. Hendry<sup>1</sup>, S. Arndt<sup>2</sup>, S.F. Henley<sup>3</sup>, J.C. Faust<sup>4, 5</sup>, A.C. Tessin<sup>6</sup>, M.A. Stevenson<sup>7, 8</sup>, G.D. Abbott<sup>7</sup>, C. März<sup>4</sup>

<sup>1</sup>University of Bristol, UK, <sup>2</sup>Université Libre de Bruxelles, Belgium, <sup>3</sup>University of Edinburgh, UK, <sup>4</sup>University of Leeds, UK,

<sup>5</sup>University of Bremen, Germany, <sup>6</sup>Kent State University, USA, <sup>7</sup>Newcastle University, UK, <sup>8</sup>Durham University, UK

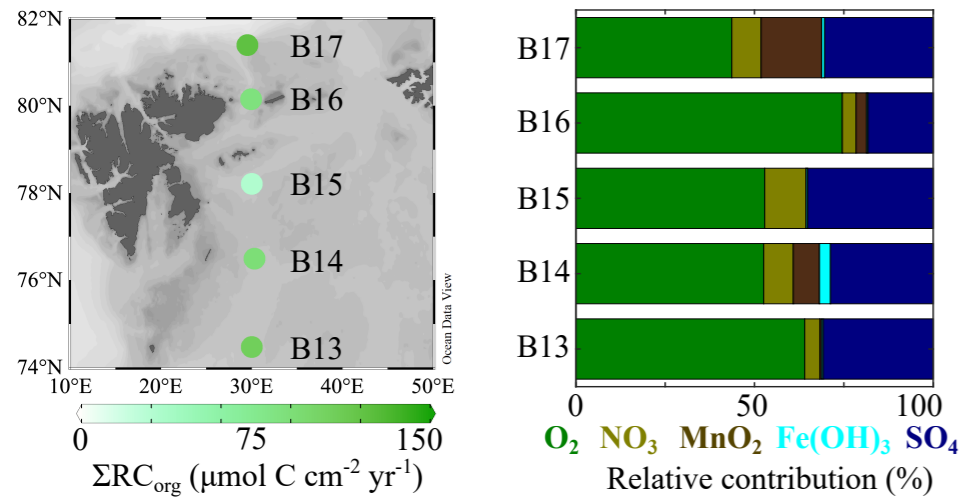
✉ felipe.salesdefreitas@bristol.ac.uk

🐦 @FelipeSdFreitas



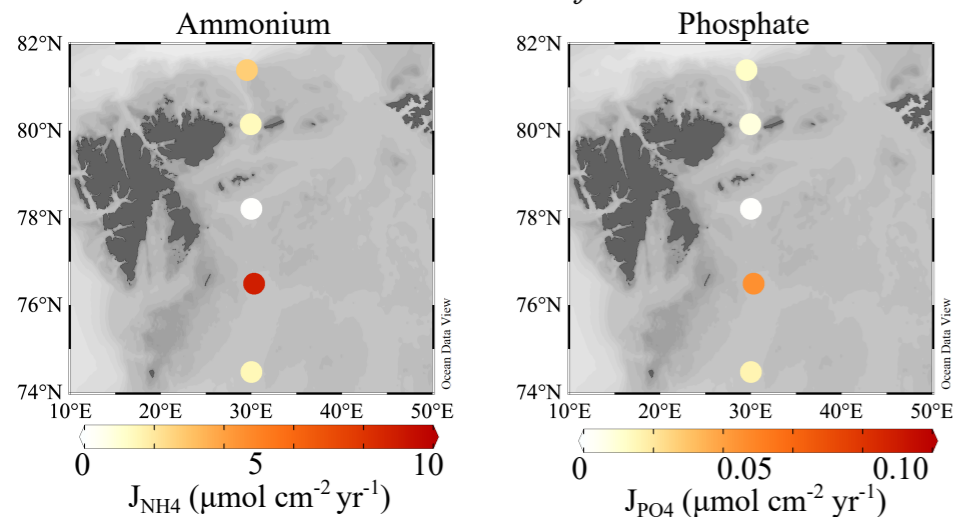
## Organic matter transformation

Rates of heterotrophic degradation



Depth-integrated (0–100 cm) rates of **OM degradation** are driven by **reactivity** patterns along the shelf. The highest rates (high reactivity) are found in **Atlantic Water** and have no clear links with **sea-ice cover**. **Aerobic degradation** represents > 50% of total rates. **Sub-oxic pathways** have small and variable relative contributions. **Sulfate reduction** contributes to around one-third of overall rates.

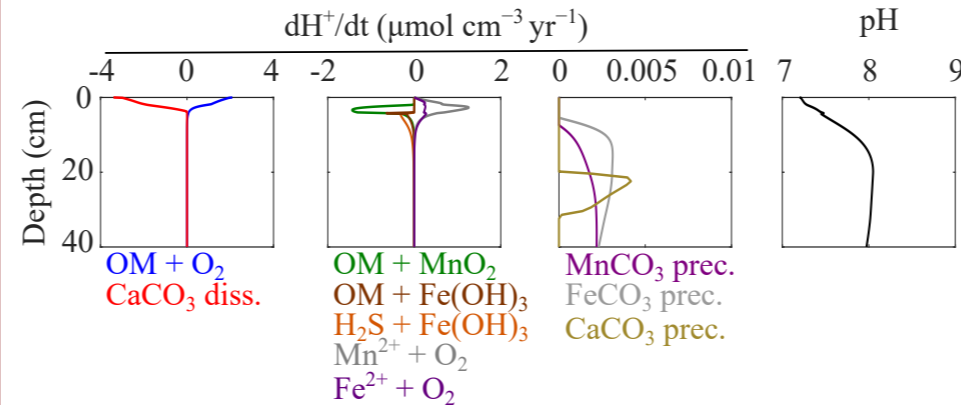
## Benthic nutrient fluxes



Seafloor efflux of **ammonium** and **phosphate** closely follow OM reactivity patterns. The lowest fluxes are found in B15 where **low degradation rates of unreactive OM** are found in association with deeper **oxygen penetration** ( $\geq 3$  cm). B14 displays the largest fluxes, where shallow **oxygen depletion** (< 0.5 cm) prevents ammonium oxidation and phosphate adsorption to iron mineral phases.

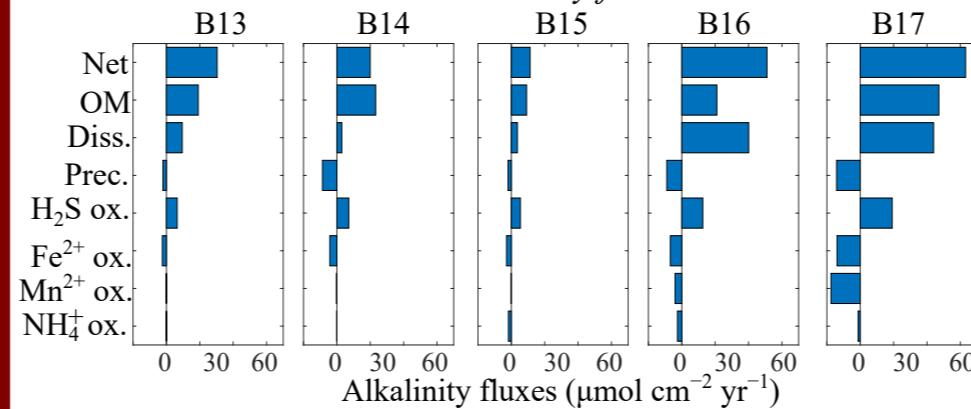
## Seafloor carbonate chemistry and pH

Biogeochemical controls on pH dynamics

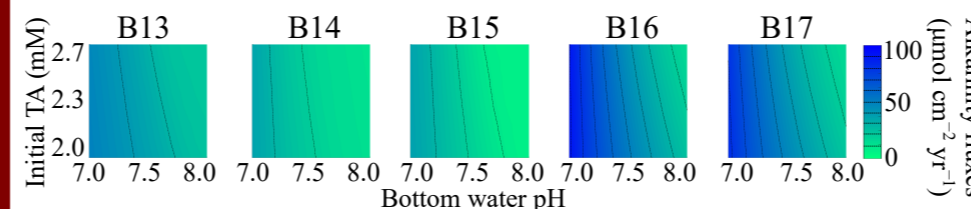


**Aerobic processes** in upper sediments (< 2 cm) produce  $\text{H}^+$  and a negative shift in pH. A reversal in pH occurs via **dissolution of  $\text{CaCO}_3$**  and **metal oxide pathways** through  $\text{H}^+$  consumption. Thus, pH increases to  $\approx 8.0$  with depth. **Carbonate saturation state** also increases, enabling authigenic precipitation.

## Benthic alkalinity fluxes



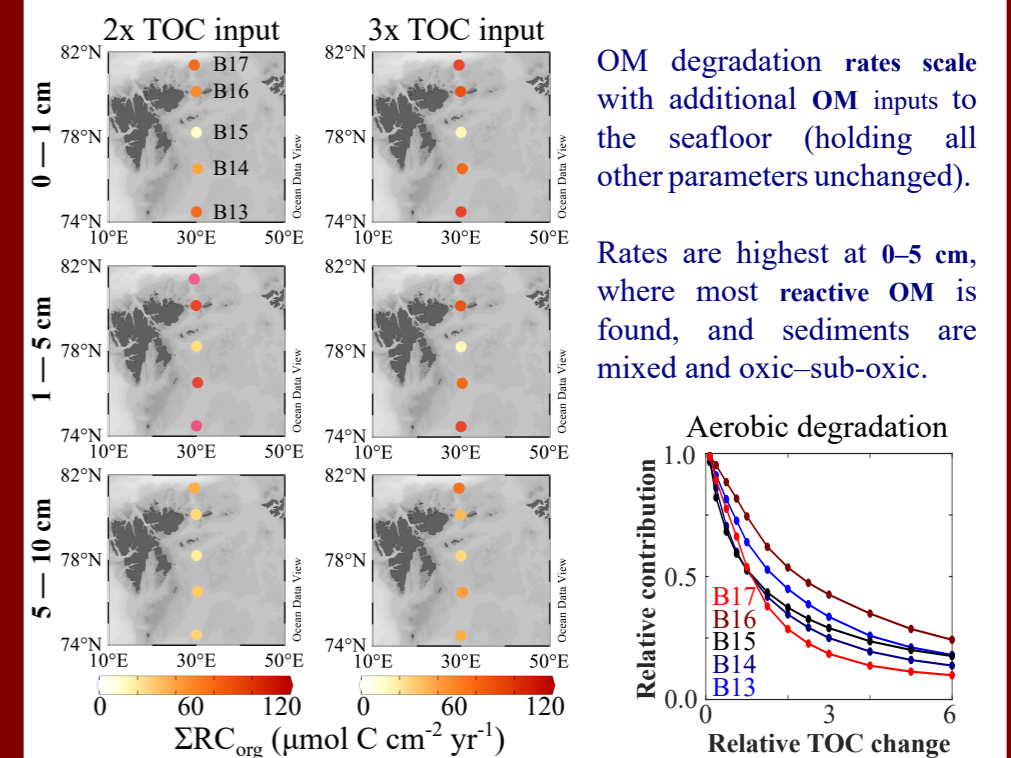
There is a **south-north** trend in alkalinity efflux. The greater fluxes in B16-B17 are fuelled by **OM degradation** and  **$\text{CaCO}_3$  dissolution**. Fluxes in B13-B15 are lower due to lower **calcite contents**, which limits dissolution rates. **Authigenic precipitation** and **oxidation processes** are the main alkalinity sinks.



Bottom water **acidification** has a large role on shallow **dissolution of calcite**, which enhances seafloor alkalinity release.

## Changes in organic matter input fluxes

Degradation dynamics

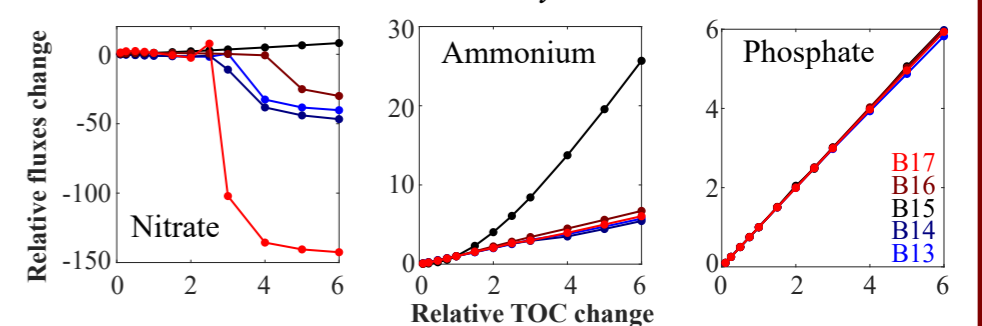


OM degradation rates scale with additional OM inputs to the seafloor (holding all other parameters unchanged).

Rates are highest at 0–5 cm, where most reactive OM is found, and sediments are mixed and oxic-sub-oxic.

Enhanced OM fluxes to the seafloor lead to **faster oxygen consumption** and a shift from **oxic to anoxic** degradation pathways.

## Nutrient dynamics



**Ammonium** and **phosphate** fluxes closely follow OM input. **Nitrate** shows a more dynamic behaviour. It changes from **release** to the bottom waters in low OM input to **uptake** in high OM degradation.

## Impacts of bottom trawling

**Fisheries** expansion in Arctic shelves brings large uncertainties to **OM storage**. Disturbance of upper sediments ( $\approx 10$  cm) may remove high reactivity OM and re-expose buried, unreactive OM, thus potentially diminishing OM burial. Nutrient fluxes are also vulnerable to **trawling**.